

# **A shallow water ocean acoustics inverse problem**

**David Stikler**

Department of Mathematical Sciences and  
Center for Applied Mathematics and Statistics  
New Jersey Institute of Technology, Newark, NJ 07102

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# A SHALLOW WATER OCEAN ACOUSTICS INVERSE PROBLEM

David Stickler

Department of Mathematical Sciences

The New Jersey Institute of Technology

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## Abstract

A shallow water ocean acoustics experiment is described from which it is possible to recover the scattering data necessary to recover the sound speed in the ocean bottom.

For this shallow water ocean acoustics inverse problem the following assumptions are made:

- The acoustic sound speed in the ocean and in the bottom depend only on the depth coordinate  $z$ ,  $c = c(z)$ , for  $0 \leq z < \infty$ .
- The density is a constant for all  $z$ .
- The ocean occupies the region  $0 \leq z \leq L$  where  $L$  is known, and the sound speed in the ocean is known. In an earlier paper[1], the sound speed in the ocean was also assumed unknown.
- The ocean bottom occupies the region  $L < z < \infty$  and the sound speed there is unknown and is to be determined.
- The sound speed approaches a constant  $c_1$  sufficiently rapidly as  $z$  approaches infinity and the constant  $c_1$  is known.

The object is to determine  $c(z)$  for  $z > L$  from measurements in the ocean  $0 \leq z \leq L$ .

The experiment to determine the data to recover  $c(z)$  is described below.

- At a depth  $z_0$  in the ocean an isotropic, low-frequency ( $\omega$  is its angular frequency) source projects acoustic energy. The frequency is chosen to be low enough so that no eigenmodes are excited. This means that only the continuous spectrum is excited.
- The pressure field is measured at the same depth for all range points that is for  $0 \leq r < \infty$ .

The Hankel transform of the pressure field is calculated .

The Hankel transform of the pressure field satisfies (the transform variable is denoted by  $\mu$ ).

(Eq 1)  $\frac{d^2 g(z, z_0, k)}{dz^2} + (k^2 - q(z))g(z, z_0, k) = \delta(z - z_0)$  for  $0 \leq z < \infty$  and  $0 < z_0 < L$ , where

(Eq 2)  $g(0, z_0, k) = 0$ .

and

(Eq 3)  $k^2 = (\omega/c_1)^2 - \mu^2$

and

(Eq 4)  $q(z) = \omega^2(1/c_1^2 - 1/c^2(z))$

Let  $S(z, k)$  and  $U(z, k)$  both satisfy the homogeneous form of Eq 1 with

(Eq 5)  $U(z, k) \sim e^{ikz}$  as  $z \rightarrow \infty$

and

(Eq 6)  $S(0, k) = 0$  and  $\frac{dS(0, k)}{dz} = 1$ .

The measured data needed to recover  $q(z)$  for  $L < z < \infty$  is [2],[3]

(Eq 7)  $S_c(k) = U(0, -k)/U(0, k)$  for  $0 \leq k < \infty$ .

For  $k$  real  $U(0, -k) = \overline{U(0, k)}$  and therefore the modulus of  $S_c(k)$  is unity.

The green's function is given by

(Eq 8)  $g(z, z_0, k) = -S(z, k)U(z_0, k)/U(0, k)$  for  $z < z_0$   
 $= -S(z_0, k)U(z, k)/U(0, k)$  for  $z_0 < z$

and since there are no eigenvalues  $U(0, k) \neq 0$ .

The measurement at  $z_0$  yields  $g(z_0, z_0, k)$  and will be denoted by  $g_M$  . In addition, since the sound speed is known in the ocean  $S(z_0, k)$  can be calculated.

From Eq 8

$$(Eq 9) U(z_0, k) = -g_M U(0, k) / S(z_0, k)$$

For real k

$$(Eq 10) S(z_0, k) = \Im(\overline{U}(0, k) U(z_0, k)) / k$$

hence

$$(Eq 11) |U(0, k)|^2 = k S^2(z_0, k) / \Im \overline{g}_M$$

But

$$(Eq 12) U(0, k) = |U(0, k)| \exp i\Theta(k).$$

However,  $\Theta(k)$  and  $\ln(|U(0, k)|)$  are the real and imaginary parts of conjugate harmonic functions and hence  $\Theta$  is given by the Cauchy principal value integral

$$(Eq 13) \Theta(k) = (-2k/\pi) \int_{k'=0}^{\infty} \ln |U(0, k')| / ((k')^2 - k^2) dk'$$

and

$$(Eq 14) S_c(k) = \exp(-2i\Theta(k)).$$

## References

- 1 D. C. Stickler, P. Deift, Inverse Scattering for a Stratified Ocean and Bottom, J. Acoust Soc. Amer. 7 (1981) pp. 1723-1727
- 2 P. Deift , E Trubowitz Inverse Scattering on the Line , Comm Pure and Applied Math. 32 (1979 ) pp 121-251
- 3 Z. S. Agranovic, V. A. Marchenko, Inverse Scattering Theory, Gordon and Breach, 1963